Carbon dioxide emission from raised bog surface after peat extraction

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Abstract

Research on CO_2 emission from a raised bog after completion of peat extraction was performed in 2011–
2013. CO_2 emissions were determined by the chamber method. Twenty years after the termination of peat ex-
traction, the bog surface was almost entirely devoid of plants. CO_2 emission from the bog varied depending on
temperature and water conditions and was 418 mg·m^{-2}·h^{-1} on average during the research period. CO_2 losses on
the raised bog were on average 19.7 Mg·ha^{-1}·year^{-1} during the research period which corresponded to a carbon
loss of 5.37 Mg·ha^{-1}·year^{-1} or mineralisation of 9.6 Mg·ha^{-1}·year^{-1} of organic mass of 56% carbon content. It is
possible to reduce organic mass losses and CO_2 emission to the atmosphere from the bog surface after peat ex-
traction has been terminated by reconstruction of initial water conditions, i.e. retaining a high ground water level
and restoration of aquatic plant communities.

Key words: *CO_2 emission, mineralization, raised bog, total ecosystem respiration*

INTRODUCTION

Peat used to be extracted from raised bogs mainly to obtain energetic material. As more caloric fossil fuels appeared on the market, significance of peat decreased considerably. In recent years, peat deposits are extracted from raised bogs mainly to produce substrates for gardening purposes. The area of exploited raised bogs in the mid-1980s was estimated at 3763 ha [LIPKA 1984], whereas the area of raised bogs under extraction in 1996 was 1062 ha [ILNICKI 2002], which is about 0.5% of the total bog area in Poland.

One of the methods of raised bogs exploitation is milling. This method consists in lowering of the ground water level in a bog, cutting thin peat layers of several dozen mm, drying and collecting them [BIE-

NIEK, ŁACHACZ 2010]. This method results in a complete destruction of plan communities. It was observed that after exploitation of a peat deposit, despite restoration attempts, post-exploitation surfaces are not settled by plants and undergo gradual degradation, mainly due to mineralisation of organic mass resources [HERBICHOWA *et al*. 2009; ROCHEFORT 2003].

Spontaneous plant regeneration on raised bogs, degraded by peat exploitation, occurs to a very limited degree, due to unfavourable habitat conditions. Furthermore, as a result of peat milling, top layer of peat bog is completely removed over large areas. Along with the diasporas from which vegetation could reproduce. The main reason for a lack of regeneration is a persistent low ground water level due to drainage of peat deposits. The lack of protective effect of water

and plants causes the surface of the peats to heat up, which hampers seed germination and seedling development [GIRARD *et al.* 2002; HERBICHOWA *et al.* 2009; SZAJDA, ŁABĘDZKI 2017]. An additional difficulty for plant development can also be extremely low pH of the substrate – even for acidophilous peat mosses occurring in raised bogs [MONEY 1994] and nutrient deficiencies.

As a result, even many years after the termination of peat exploitation, post-extraction sites remain almost entirely devoid of vegetation. Lack of vegetation results in a deterioration of CO_2 balance and increased carbon losses. In Poland there has been no research into carbon losses on raised bogs after peat extraction termination. The aim of the paper is to determine CO_2 emissions and carbon losses on a raised bog after termination of peat extraction.

RESEARCH OBJECT AND METHODS

The research was conducted on Czarne Bagno raised bog situated in the Łeba valley (Fig. 1) at a distance of about 10 km west from Lębork (Pomeranian province). Czarne Bagno is a Baltic raised bog. The original area of the bog dome was 226 ha, and peat thickness reached 8 m, of which about 1.5 m was raised moss peat. At the end of the 19th century Czarne Bagno was included in the system of ditches draining the whole Łeba ice marginal streamway. Particularly intensive drainage was carried out about 1970, which resulted in lowering of water table by 0.5 m in the whole area. Till the end of the 1950s the bog was covered by extensive peat extraction, and in 1987–1989 peat was extracted on an industrial scale by the milling method in an area of 9.2 ha. As a result, about one-metre-thick peat layer was harvested. In

Fig. 1. Location of research area: *1* = CO_2 emission measuring site, *2* = piezometer, *3* = meteorological station, *4* = main ditches, *5* = nature reserve boundary, *6* = milling extraction areas; source: own elaboration

2006 a nature reserve was established on an area of 102.86 ha and renaturation steps were taken [DUDA *et al.* 2016; GDOŚ 2015, HERBICHOWA *et al.* 2007].

CO_2 emission measurements were performed in 2011–2013, from April to October, at decade intervals by the chamber method, repeated twice. CO_2 concentration changes in the chamber were determined using a photoacoustic meter (once a month) by taking a gas sample from the chamber to the meter in a closed circuit, and a diffusion meter (once-twice a month), which was placed inside the chamber. Dimensions of the chamber were 45 × 45 × 35 cm. It was made of transparent acrylic glass and equipped with a ventilator, in order to maintain a uniform gas concentration inside, and a valve to equalise pressure between the chamber and the atmosphere. The chamber was sequentially placed in two square channel-bar frames permanently fixed to the analysed surface. The space between the frame and the chamber was sealed with water filling the frame. The bottom of the frame had blades which were driven into the ground surface to a depth of 5 cm. During measurement the chamber was protected by a light-proof cover in order to reduce air temperature increase inside. Measurements were conducted between 10:30 am and 11:30 am. During measurement the air temperature inside the chamber was measured. CO_2 emission measurement lasted about 6–8 minutes and the frequency of recording – every minute. In order to determine CO_2 fluxes values the measurement of the first 2–3 minutes was used when CO_2 concentration increase was linear. In 2011, 2012 and 2013 respectively, 16, 20 and 17 measurements of CO_2 emissions were made, a total of 53. CO_2 concentration changes in ppm were converted into $mg \cdot m^{-2} \cdot h^{-1}$, according to the following formula [MOSIER, MACK 1980]:

$$E = \rho \cdot V/A \cdot \Delta C/\Delta t \cdot 273/(T + 273) \qquad (1)$$

where: E = flux value, $mg \cdot m^{-2} \cdot h^{-1}$; ρ = gas density, $mg \cdot m^{-3}$; V = chamber capacity, m^3; A = chamber area, m^2; $\Delta C/\Delta t$ = mean rate of gas concentration changes in time, $ppmv \cdot h^{-1}$; T = temperature inside chamber, °C.

CO_2 emission annual value was calculated using the regression equation describing the relationship between soil respiratory activity and air temperature according to LLOYD and TAYLOR [1994]:

$$TER = R_{ref} \cdot \exp\{E_0 \cdot [(1/(T_{ref} - T_0)] - [1/(T - T_0)]\} \quad (2)$$

where: TER = total respiratory activity, $mg \cdot m^{-2} \cdot h^{-1}$; R_{ref} = respiration in reference temperature at 10°C, $mg \cdot m^{-2} \cdot h^{-1}$; E_0 = activation energy coefficient; T_{ref} = reference temperature – 283.15 K; T_0 = constant temperature in which biological processes are initiated – 227.1 K; T = soil or air temperature, K.

Function parameters R_{ref} and E_0 were matched to the data set using Statistica 7.1 programme. The obtained regression equation was used for modelling the
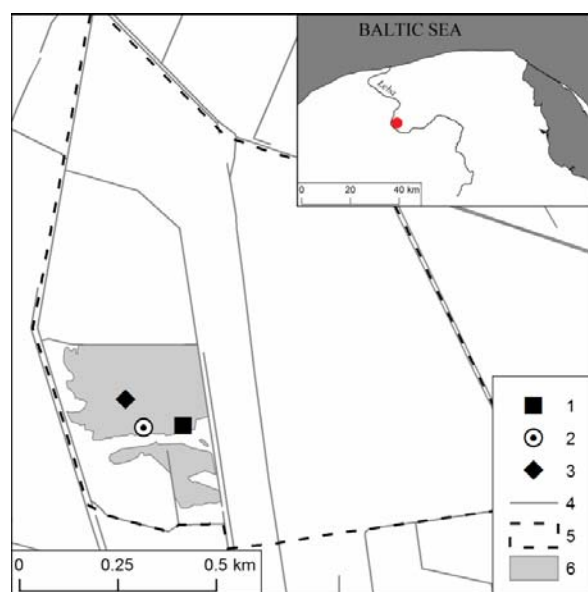
Table 1. Mean monthly air temperature, °C

Year	Temperature in month												Mean
	I	II	III	IV	V	VI	VII	VIII	IX	X	XI	XII	
2011	−0.2	−3.9	3.1	10.1	13.0	17.8	18.2	17.6	14.0	8.8	5.1	3.2	8.9
2012	0.1	−4.1	3.8	7.2	12.4	14.4	17.6	16.9	13.6	7.8	5.3	−1.4	7.8
2013	−1.5	−0.2	−1.7	6.2	13.6	15.7	17.5	16.7	11.5	9.7	5.4	2.8	8.0
Mean	−0.5	−2.7	1.7	7.8	13.0	16.0	17.8	17.1	13.0	8.8	5.3	1.5	8.2

Source: own elaboration.

Table 2. Mean ground water levels, cm

Year	Ground water level in month												Mean
	I	II	III	IV	V	VI	VII	VIII	IX	X	XI	XII	
2011	−13.7	−17.5	−19.1	−19.4	−32.2	−43.0	−39.0	−33.5	−25.6	−23.2	−22.0	−13.3	−25.1
2012	−1.5	0.0	0.0	−0.6	−10.1	−19.5	−5.7	−1.1	0.0	0.0	0.0	0.0	−3.2
2013	0.0	0.0	0.0	0.0	0.0	−11.5	−21.8	−23.8	−20.7	−17.8	−11.6	−6.6	−9.5
Mean	−0.1	−0.6	−1.6	−4.6	−14.0	−24.7	−22.2	−19.5	−15.1	−10.9	−6.8	0.6	−10.0

Source: own elaboration.

course of respiratory activity in the following years, using air temperature measurements taken hourly. Air temperature measurements were performed on an automatic meteorological station situated within the research object (Tab. 1).

The data for ground water level in the bog area where milling extraction was performed was obtained from the hydrological monitoring network. Mean daily water levels (for measurement days) from a piezometer situated 100 m from the CO_2 emission measurement site were used. The average water levels in 2011–2013 are presented in Table 2. The coordinates were obtained from the Digital Terrain Model [CODGiK 2011].

Ash content of soil samples was determined by ashing at a temperature of 550°C, soil pH in 1 N KCl. Water properties of soil were determined by ZAWA-DZKI [1973] method by taking soil into cylinders of 100 cm^3 capacity in four repetitions from subsequent 10-cm-thick layers.

RESULTS

CO_2 emission measurements were performed on the surface of a bog after industrial peat extraction was terminated. Even twenty years after the termination of peat extraction, the surface of the soil where CO_2 emission was measured was almost entirely devoid of plants, mainly due to very low pH values (Tab. 3).

In 2011–2013 from April to October the mean CO_2 emission value was 418 mg·m^{-2}·h^{-1}. The highest CO_2 emission values were recorded in 2011 – on av-

erage 469 mg·m^{-2}·h^{-1}, and the lowest in 2013 – on average 365 mg·m^{-2}·h^{-1} (Tab. 4). The higher emission value in 2011 than in 2012–2013 was caused by persistently low ground water level of 25.1 cm below ground level. Favourable oxygen conditions stimulated the respiratory activity of soil microorganisms thus increasing CO_2 emission.

In particular months of the research period the lowest value was recorded in April – on average 176 mg·m^{-2}·h^{-1}. In May the mean CO_2 emission value was almost three times higher than in April and was 497 mg·m^{-2}·h^{-1}. The highest emission value was recorded in June – 690 mg·m^{-2}·h^{-1}. In subsequent months i.e. July, August and September the emission values diminished reaching 300 mg·m^{-2}·h^{-1} in October (Tab. 4). The lowest CO_2 emission values were observed in the months with the lowest air temperatures i.e. in April and October (7.8 and 8.8°C – Table 1).

No close correlation was however observed between air temperature and CO_2 emission in summer months. Although the highest mean temperature was recorded in July (17.8°C). CO_2 emission in this month was 514 mg·m^{-2}·h^{-1} and was by 25.5% lower than in June when mean temperature was 16.0°C. This indicates that apart from temperature. CO_2 emission value is considerably influenced by water conditions.

An analysis of the relationship between respiratory activity and ground water table level indicates a very strong diversity of CO_2 emissions with the same ground water level (Fig. 2). Even in conditions of a complete saturation of the soil profile with water very high CO_2 emission values were observed. When ground water level remained at the soil surface. CO_2

Table 3. Soil physical and water properties

Layer	Ash content, % DM	Bulk density, Mg·m^{-3}	pH	Water capacity (cm^3·cm^{-3}·100) at pF			
	$\overline{x} \pm SD$	$\overline{x} \pm SD$	1 N KCl	0.0	2.0	2.7	4.2
0–10	7.51±0.37	0.094 ± 0.009	2.52	98.4	79.4	45.9	11.1
10–20	6.11±0.25	0.097 ± 0.003	2.28	98.4	78.0	43.3	11.9
20–30	2.67±0.15	0.078 ± 0.004	2.30	98.6	79.5	35.6	12.5
30–40	3.92±0.16	0.073 ± 0.003	2.36	98.6	76.3	32.0	8.7

Explanations: $\overline{x} =$ mean value, $SD =$ standard deviation.
Source: own study.

Table 4. CO_2 emission from bog surface after peat extraction

Month	CO_2 emission a year, $mg \cdot m^{-2} \cdot h^{-1}$			Mean $\bar{x} \pm SD$
	2011	2012	2013	
IV	222 ± 110*	255 ± 38	51 ± 35	176 ± 110
V	878 ± 134	338 ± 209	277 ± 120	497 ± 331
VI	834 ± 15	699 ± 38	537 ± 168	690 ± 149
VII	493 ± 288	665 ± 63	382 ± 39	514 ±142
VIII	324 ± 85	428 ± 220	513 ± 170	422 ± 95
IX	272 ± 162	281 ± 102	431 ± 291	328 ± 89
X	264 ± 104	271 ± 147	365 ± 83	300 ± 57
Mean	469 ± 276	420 ± 214	365 ± 192	418 ± 52

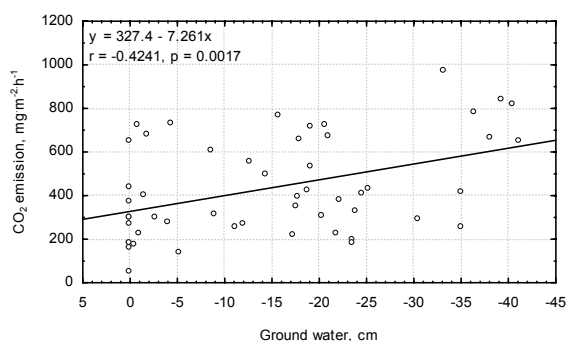Explanations as in Tab. 3.
Source: own study.

Fig. 2. Relationship between CO_2 emission and ground water level; source: own study

emission was from 51 $mg \cdot m^{-2} \cdot h^{-1}$ on 08.04.2013 to 722 $mg \cdot m^{-2} \cdot h^{-1}$ on 04.06.2012. The reason for the persistence of high soil respiratory activity in conditions of long lasting high ground water level was the ability of microorganisms to use oxygen from oxidised mineral compounds such as iron and manganese oxides as proton acceptor [GLIŃSKI *et al.* 1983; STĘPNIEWSKA *et al.* 2004]. What follows is that even a short lowering of ground water level resulting in oxidation of mineral compounds makes the mineralisation process possible to proceed in condition of a complete filling of soil pores with water.

A high diversity of CO_2 emission was also observed when ground water level remained at a depth of about 35 cm below ground level. In the first half of 2011 (from 12.05 to 14.07). CO_2 emission was about 700 $mg \cdot m^{-2} \cdot h^{-1}$, whereas in the second half of this year – about 300 $mg \cdot m^{-2} \cdot h^{-1}$ (Fig. 3). Such a substantial decrease in CO_2 emission was partially related to lower air temperature in that period and could also be related to a depletion of the supply of easily decomposed organic compounds by aerobic microorganisms or a termination of their intense development which resulted in diminished respiratory activity.

The most favourable conditions in terms of reducing CO_2 emission from post-exploitation areas were those when ground water level remained high for a long time. In the period from the end of August 2012 to the end of May 2013, when ground water level remained close to soil surface, a clear reduction in CO_2 emission value was observed. In that period only one record revealed emission above 400 $mg \cdot m^{-2} \cdot h^{-1}$, and on 26.05.2013 it was only 161 $mg \cdot m^{-2} \cdot h^{-1}$. A lowering of ground water level in mid-June 2013 due to high air temperature did not however cause a rapid CO_2 emission increase. Throughout July CO_2 emission remained at a level of about 400 $mg \cdot m^{-2} \cdot h^{-1}$, and its gradual increase was recorded only in the second half of August (Fig. 3). As a result CO_2 emission value was the smallest in 2013 at a level of 365 $mg \cdot m^{-2} \cdot h^{-1}$.

It was observed that the strongest reduction of the organic mass mineralisation rate in peat soils can be obtained only in conditions of long-lasting high ground water level. Changes in these conditions even a short-term lowering of ground water level resulted in an increase in CO_2 emission values.

The mean values of CO_2 emission from the post-extraction surface ranged within the scope of CO_2 emission variability determined for low bogs entirely devoid of plants and retained as bare fallow. On low bogs in the Noteć valley in four sites differing in water conditions the highest respiratory activity was observed in moist and dry complexes – 510 and 470 $mg \cdot m^{-2} \cdot h^{-1}$ CO_2 respectively then in a periodically dry one – 420 $mg \cdot m^{-2} \cdot h^{-1}$ whereas the lowest in a wet complex – 330 $mg \cdot m^{-2} \cdot h^{-1}$ [TURBIAK *et al.* 2009].
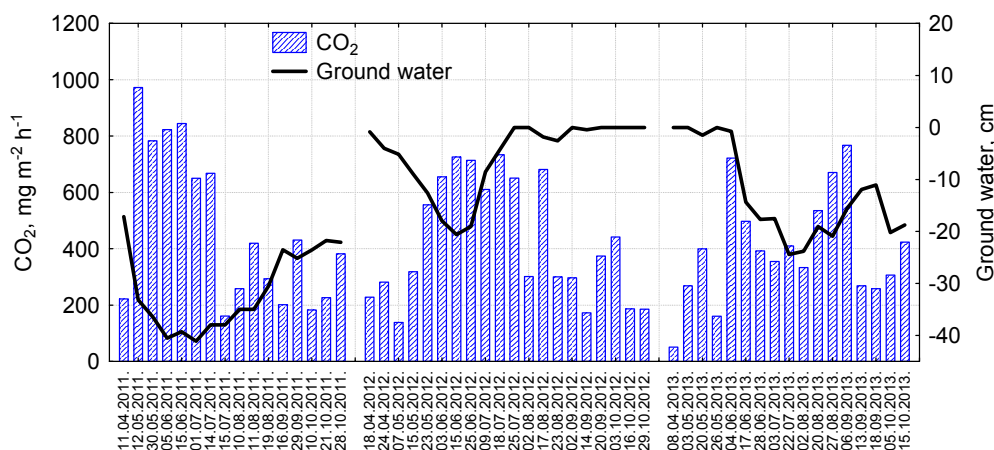

Fig. 3. CO_2 emission from bog surface and ground water level in 2011–2013; source: own study

Higher values were obtained by AERTS and LUDWIG [1997] who determined CO_2 emission from eutrophic and mesotrophic bogs at 13.1 and 15.2 $g \cdot m^{-2} \cdot d^{-1}$ respectively i.e. 546 and 633 $mg \cdot m^{-2} \cdot h^{-1}$. Similar CO_2 emission figures from optimally water-logged peat-muck soils remaining in the condition of bare fallow were obtained by MALJANEN *et al.* [2004], who determined CO_2 emission at a level of 158 mg $C-CO_2 \cdot m^{-2} \cdot h^{-1}$ (580 mg $CO_2 \cdot m^{-2} \cdot h^{-1}$).

CO_2 emission value in a year was calculated on the basis of the regression equation describing the relationship between CO_2 emission and air temperature gauged during measurements (Fig. 4). Thanks to that daily and seasonal fluctuations in temperature were taken into account in calculations.
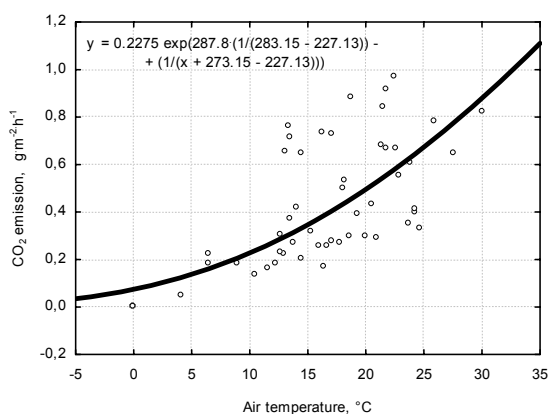

Fig. 4. Relationship between air temperature and CO_2 emission; source: own study

In 2011–2013 on average 19.7 Mg $CO_2 \cdot ha^{-1} \cdot year^{-1}$ was emitted from the surface of the raised bog after peat extraction. This accounts for a carbon loss of 5.37 $Mg \cdot ha^{-1} \cdot year^{-1}$ or 9.6 $Mg \cdot ha^{-1} \cdot year^{-1}$ loss of organic mass of 56% carbon content (Tab. 5).

Table 5. CO_2 emission value and losses of carbon (C) and of organic mass

Parameter	Losses, Mg·ha⁻¹·year⁻¹			Mean
	2011	2012	2013	
CO_2	21.8	18.3	19.0	19.7
C	5.94	4.99	5.18	5.37
Organic mass	10.6	8.9	9.2	9.6

Source: own study.

Lower carbon loses were found in closed peat mines in Estonia. In five mines the average C-CO_2 emission from the post-mining areas was 2,845 $Mg \cdot ha^{-1} \cdot year^{-1}$ [SALM *et al.* 2012].

It could be assumed that due to very low pH and small nutrient content in raised bog respiratory activity and organic mass mineralisation rate would be substantially smaller than in low bog rich in nutrients. The obtained results reveal that on a drained raised bog there is a very intensive process of organic mass mineralisation and organic mass losses are similar as in eutrophic low bogs.

The calculated carbon losses on the raised bog were almost identical as the values given by the Intergovernmental Panel on Climate Change. According to IPCC in moderate climate carbon losses on drained nutrient-poor organic soils are 5.3 $Mg \cdot ha^{-1} \cdot year^{-1}$ whereas in nutrient-rich soils they are 6.1 $Mg \cdot ha^{-1} \cdot year^{-1}$ [IPCC 2013].

CONCLUSIONS

1. Mean value of CO_2 emission from the bog surface after extraction was 418 $mg \cdot m^{-2} \cdot h^{-1}$ for the research period. CO_2 emission depended on temperature and water conditions. The highest emission values were observed in June while the lowest in April and October. i.e. in months with the lowest air temperatures.

2. On the raised bog after peat exploitation there is a very intensive process of organic mass mineralisation at a level similar to that in low eutrophic bogs. CO_2 losses on the raised bog in the research period were 19.7 $Mg \cdot ha^{-1} \cdot year^{-1}$ on average which corresponded to a carbon loss of 5.37 $Mg \cdot ha^{-1} \cdot year^{-1}$ or mineralisation of 9.6 $Mg \cdot ha^{-1} \cdot year^{-1}$ of organic mass of 56% carbon content.

3. It is possible to reduce organic mass losses and CO_2 emission to the atmosphere from the bog surface after peat extraction by reconstruction of initial water conditions i.e. retaining a constant and high ground water level and restoration of plant communities.

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Emisja ditlenku węgla z powierzchni torfowiska wysokiego po zakończonej eksploatacji torfu

STRESZCZENIE

Badania emisji CO_2 z torfowiska wysokiego po zakończonej eksploatacji torfu prowadzono w latach 2011–2013. Emisję CO_2 oznaczano metodą komorową. Powierzchnia torfowiska po dwudziestu latach od zakończenia eksploatacji torfu była prawie całkowicie pozbawiona roślin. Emisja CO_2 z torfowiska była zależna od temperatury i warunków wodnych. Wynosiła ona średnio w okresie badań 418 mg·m⁻²·h⁻¹. Straty CO_2 na torfowisku wysokim wynosiły średnio w okresie badań 19,7 Mg·ha⁻¹·rok⁻¹, co związane było z ubytkiem 5,37 Mg·ha⁻¹·rok⁻¹ węgla lub mineralizacją 9,6 Mg·ha⁻¹·rok⁻¹ masy organicznej o zawartości 56% węgla. Ograniczenie strat masy organicznej i emisji CO_2 do atmosfery z powierzchni torfowiska po zakończonej eksploatacji torfu jest możliwe po odtworzeniu pierwotnych warunków wodnych, polegających na utrzymywaniu wysokiego poziomu wody gruntowej oraz restytucji zbiorowisk roślinnych.

Słowa kluczowe: *całkowita respiracja ekosystemu, emisja CO_2, mineralizacja, torfowisko wysokie*